



Experimental observations of heterogeneous strains inside a dual porosity sample under the influence of gas-sorption: A case study of fractured coal

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ABSTRACT

A “permeability equilibration time” is typically assumed in interpreting permeability measurements – indicating that equilibration has been reached and both sorption-induced changes in deformation and their impact on permeability evolution have ceased. However, for extremely low matrix permeability (tight) dual porosity rocks, this “permeability equilibration time” may easily exceed the time interval between two consecutive permeability measurements – invalidating the interpretation of a steady permeability if the non-steady state conditions are not correctly accommodated. This is especially important where pressure diffusion from fracture to matrix results in a non-monotonic and non-asymptotic approach to a steady permeability, but instead contains multiple stages, plateaus and permeability reversals. We validated this hypothesis through experiments and analysis. Experiments measured the non-monotonic and scale-dependent deformations of fracture and matrix and linked these directly to the dynamic evolution of reservoir permeability. These laboratory strain measurements were integrated with numerical analyses to explore how mass and stresses transferred between matrix and fracture and were coupled under conditions of constant confining pressure. Strain gauges were distributed to directly measure stress transfer between matrix and fracture and interrogated deformation at different scales and at different proximities to control fractures. The prismatic sample of coal was tested under freely expanding boundary conditions. Optical microscopy and X-ray CT imaging were used to define the fracture distribution throughout the sample with mercury intrusion (capillary) porosimetry (MICP) constraining the pore size distribution and enabling independent estimation of matrix permeability. A numerical model was built and verified by matching measured strains and then applying this to model the evolution coal permeability from initial to ultimate equilibrium. Both the experimental and numerical results show that the final equilibrium state (pressure, stress and mass contents) for the matrix system extends to months rather than hours and suggests that some current permeability data may therefore reflect a non-equilibrium permeability state. Results also show that during this non-equilibrium condition, the swelling of the matrix near the fracture will cause not only compaction and narrowing of the fracture, but also shrinkage of the matrix that is distant from the fracture under constant confining pressure condition. Both experimental and numerical results demonstrate that the evolution of non-equilibrium strain/permeability is determined by the matrix-fracture interactions, including sorption-induced swelling/shrinking, through transient stresses in matrix and fractures. And that these non-equilibrium stress transfers determine the dynamic permeability evolution during gas extraction (e.g., CH₄) or injection (e.g., CO₂) at reservoir scale for tight dual porosity rocks (e.g., coal and shale).

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List of Symbols

| | |
|------------|---|
| b_0 | Initial fracture aperture (m) |
| Δb | Fracture aperture change (m) |
| f_i | The component of body force in the i -direction (N) |
| G | Shear modulus (MPa) |
| E | Young's modulus (MPa) |
| K | Bulk modulus (MPa) |
| k_m | The permeability of matrix system (m^2) |
| k_f | The permeability of fracture system (m^2) |
| k_{f0} | The initial fracture permeability (m^2) |
| M | The molecular mass of gas (kg/mol) |
| p_m | The gas pressure in the matrix systems (MPa) |
| p_f | The gas pressure in the fracture systems (MPa) |
| p_L | Langmuir pressure (MPa) |

| | |
|-----------------|---|
| p | Pore pressure (MPa) |
| R | The universal gas constant (J/(mol·K)) |
| T | The absolute gas temperature (K) |
| u_i | The component of displacement in the i -direction (m) |
| V_L | Langmuir volume constant (m^3/kg) |
| Q_S | Gas source or sink ($kg/(m^3 \cdot s)$) |
| ϕ_m | The porosity of matrix system |
| ε_L | Langmuir volumetric strain constant |
| ν | Poisson's ratio |
| α | Biot's coefficient |
| ρ_s | Coal density (kg/m^3) |
| ρ_a | Gas density at atmospheric pressure (kg/m^3) |
| ρ_g | Gas density (kg/m^3) |
| μ | Dynamic viscosity of the gas (mPa·s) |

1. Introduction

The boom in unconventional resources (e.g., coalbed methane, shale gas) has substantially reshaped the oil and gas industry over the past two decades. Exploitation of coalbed methane (CBM) from coal seams also benefits mining safety and reduces greenhouse gas emissions (Karacan et al., 2011). Coal and shale reservoirs are generally regarded as dual porosity media that consists of the porous matrix and the surrounding fractures. However, gas adsorption/desorption induces swelling/shrinking of the matrix system (Karacan, 2007; Kiyama et al., 2011; Pan and Connell, 2007), further affecting the permeability evolution. The contrasting permeabilities of the adjoining matrix and fracture systems sustain a pressure difference between the matrix system and fracture system, together with a differential effective stress during non-equilibrium periods of gas extraction/injection. Meanwhile, interaction between matrix and fracture determines the dynamic change of reservoir permeability in dual porosity rocks and exerts a control on gas production during the drainage process (Cui et al., 2018a; Liu et al., 2017; Wang et al., 2016; Wei et al., 2019c). Thus, it is vital to understand modes of stress transfer between matrix and fracture in response to gas flow in dual porosity rocks for geoeengineering activities (e.g., coalbed methane production, shale gas extraction).

Coal reservoirs can be represented as a dual porosity/permeability system. The fractures are the primary pathway for rapid fluid flow, while the coal matrix serves as a storage site with gas stored in the various-sized pores (Haenel, 1992; Zheng et al., 2018). It is commonly assumed that Darcy flow predominates in the fracture system but can be neglected in the coal matrix where diffusion dominates (Cui et al., 2018b; Fan et al., 2019; Liu et al., 2019; Palmer, 2009; Pan and Connell, 2012; Purl et al., 1991; Ried et al., 1992; Zhang et al., 2019) – although there are equivalencies between permeability and diffusivity. When gas is injected into coal, gas rapidly invades the fractures due to the relatively high permeability. As a consequence, a pressure difference between the matrix and the fracture is created and this in turn results in the diffusion of gas from the fracture into the matrix. The gas diffuses rapidly from the fractures into the matrix where there is a high matrix permeability/diffusivity (e.g., sandstone) – and the interval time for pressures to stabilize can be neglected (Shi et al., 2018). However, for coal, this may take several days (Harpalani and Chen, 1997) or even several weeks (Gensterblum et al., 2014). When, in addition to diffusion-only, the retarding impact of gas sorption is included, the time for equilibration may extend to a few months (Danesh et al., 2017; Guo et al., 2007). These coupled diffusion/sorption phenomena determine the localized deformation within the ensemble matrix-fracture system and can be investigated through tightly constrained laboratory experiments and numerical models.

Single porosity/permeability models (Connell et al., 2010; Cui and Bustin, 2005; McKee et al., 1988; Palmer and Mansoori, 1996; Pan and

Connell, 2007; Pini et al., 2009; Robertson and Christiansen, 2006; Seidle et al., 1992; Seidle and Huitt, 1995; Shi and Durucan, 2004) are used under specific conditions (such as uniaxial strain) to explain the experimental observations. However, these models are not applicable in interpreting some experimental observations (Izadi et al., 2011; Liu et al., 2011b; Wang et al., 2011). On the basis of dual poroelasticity theory, dual porosity and dual permeability models have been established (Bai et al., 1993; Lu and Connell, 2007, 2011; Pan and Connell, 2007; Peng et al., 2014; Wang et al., 2013; Wu et al., 2011; Wu et al., 2010), which are capable of accommodating the roles of fracture-matrix interaction. Among these models, the interactions between coal matrix and fractures are normally defined by the mass exchange of gas. However, the role of mechanical interaction between matrix and fracture, that can cause the transition of coal matrix swelling, from local swelling to macro-swelling, under differential pressure, together with their impact on the evolution of permeability is not rigorously accommodated (Liu et al., 2011a; Liu et al., 2011b). To address this shortcoming, permeability models have been established with the interaction between matrix and fracture geometry and location suitably accommodated (Liu et al., 2018; Wei et al., 2019a; Zhang et al., 2018). In this, a full set of cross-coupling relations are connected between the matrix and the fracture, including local force balance, local deformation compatibility, and mass exchange (Zhang et al., 2018). The fracture aperture change in coal containing discrete fractures, following gas injection under unconstrained conditions, is consequently rigorously accommodated (Liu et al., 2018). This includes consideration of the non-uniform deformation induced by gas diffusion and described by a strain-rate-based permeability model coupling coal deformation and gas flows in both fractures and matrix (Wei et al., 2019a).

The majority of experimental studies on cleat-matrix interaction

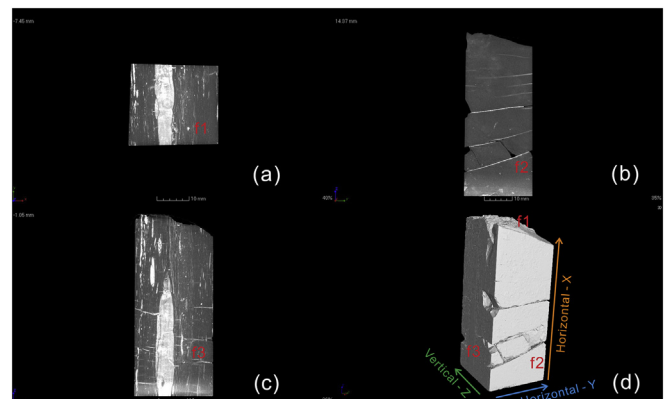


Fig. 1. CT images of the coal sample for faces: (a) f1, (b) f2, and (c) f3, (d) together with a three-dimensional view of the coal sample.

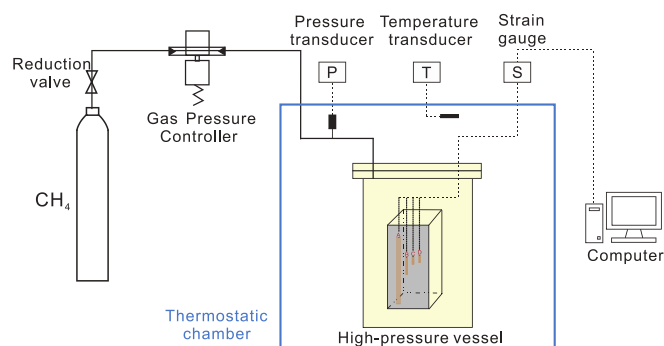


Fig. 2. Schematic of the experimental set-up to measure the sorption-induced strain.

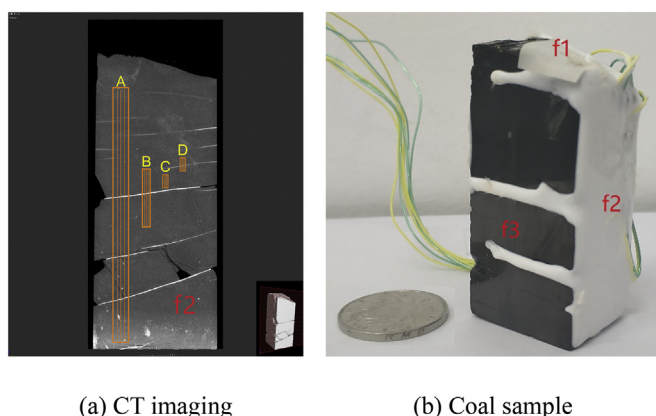


Fig. 3. Schematic and photograph of the sample with strain gauge arrangement.

have been completed by measuring the correlation between coal permeability and pore pressure – with these experiments divided between displacement-controlled and stress-controlled experiments. For displacement-controlled experiments, uniaxial strain experiments are normally used to study the evolution of coal permeability (Fan and Liu, 2018; Mitra et al., 2012; Wang et al., 2015). As for the stress-controlled experiments, two types of experiments are generally conducted, with one keeping total stress constant (Gensterblum et al., 2014; Harpalani and Schraufnagel, 1989, 1990; Harpalani and Zhao, 1989; Kumar et al., 2012; Meng et al., 2015; Pini et al., 2009; Robertson and Christiansen, 2005; Wang et al., 2017b; Wang et al., 2019; Wang et al., 2015), and the other keeping effective stress constant (Al-hawaree, 1999; Anggara et al., 2016; Chen et al., 2011; Feng et al., 2017; Harpalani and Chen, 1997; Li et al., 2015; Li et al., 2013; Lin and Kovscek, 2014; Lin et al., 2008; Meng and Li, 2017; Pan et al., 2010; Seomoon et al., 2015; Xu et al., 2013). Regardless of the choice of boundary conditions, these experiments typically ignore the pressure difference between the matrix and the fracture – and assumes steady condition to interpret the results. In order to study the impact of dynamic deformation between matrix and fracture within the coal sample, a set of experiments have been performed to measure the evolution of dynamic permeability during gas injection/depletion under constant confining pressures (Liu et al., 2016a; Mazumder and Wolf, 2008; Siriwardane et al., 2009; Wei et al., 2019b). Importantly, the evolution of the *distribution* of strains throughout the entire samples and under different stress conditions were also observed. Helium gas (He) was injected into both jacketed and unjacketed coal samples within a pressurized core holder to investigate coal matrix swelling during the process of gas diffusion from cleats, and the impact on the evolution of strain within the entire sample were observed (Wang et al., 2016). For a constant volume boundary condition (zero displacement), local deformation of the coal

sample and its permeability were measured through strain gauges and unsteady flow during helium gas injection (Wang et al., 2017a), respectively. These results demonstrated that gas diffusion from the fracture to the matrix can result in localized swelling of the coal matrix and affect the aperture of the cleat. Under this condition, coal permeability is controlled primarily by the local deformation. Normally it is assumed that 100% of the coal swelling would contribute to the reduction of coal permeability provided that the fractures are much more compliant than the coal matrix (Harpalani and Chen, 1995; Liu et al., 2011b; Ma et al., 2011). However, few direct observations of this have been made, and the distribution of deformations in different parts of the sample has not been clearly explained.

The following study observed the stress transfer between matrix and fracture through the proxy of measured strains. Strain gauges were used to measure the mean strain on different parts of a prismatic coal sample containing a variable density distribution of fractures. The sample was tested under conditions of free expansion (zero stress). Optical microscopy and X-ray CT imaging were used to define the fracture distribution throughout the sample with mercury intrusion (capillary) porosimetry (MICP) constraining the pore size distribution measured and enabling independent estimation of matrix permeability. This work offers a first direct observation into the dynamics of stress transfer between fracture and matrix and a new understanding of permeability evolution in response to the transition in dual porosity rocks. These results and findings are reported in the following sections.

2. Experimental method

Methane was sorbed into a freely swelling prismatic sample of coal sample to measure the dynamic evolution of strain at the decimeter scale. Optical microscopy and X-ray CT imaging were utilized to define the fracture distribution throughout the sample with pre-sorption porosity determined by mercury injection capillary pressure porosimetry (MICP). Surface mounted strain gauges were attached to the sample to monitor the local linear strains of different parts of the sample and at different scales, both adjacent to fractures and in the intact portion of the sample. The sample preparation and experimental techniques are introduced in the following.

2.1. Sample information

Samples of anthracite from the Qinshui Basin were used in the experiments with basic information obtained from the Optical microscopy, X-ray CT imaging and MICP tests.

2.1.1. Sample preparation

Prismatic blocks of anthracite for the deformation tests were collected from an underground coal mine in Permian-Carboniferous strata of the Qinshui Basin, China. A single prismatic sample was cut from the coal blocks, as shown in Fig. 1 (d). The height of the sample is ~5 cm with length and width of ~2 cm. The sample was dried at 50 °C for 48 h to remove moisture.

2.1.2. Optical microscopy and X-ray CT imaging

The fracture distribution on the polished faces was recorded under reflected light using an optical microscopy (Nikon, manufactured in Tokyo, Japan). X-ray computed tomography (CT) was utilized as a non-destructive technique to provide quantitative detection of the interior 3D structure of the sample (Wang et al., 2019; Wildenschild and Sheppard, 2013). In this study, the sample was scanned with a Nanotom X-ray Computed Tomography instrument (GE Phoenix) with a resolution of 1 μm.

2.1.3. Mercury injection capillary pressure porosimetry (MICP)

The geometric characteristic of the porosity distribution and connectedness determines the diffusion coefficient of the coal (Yang and

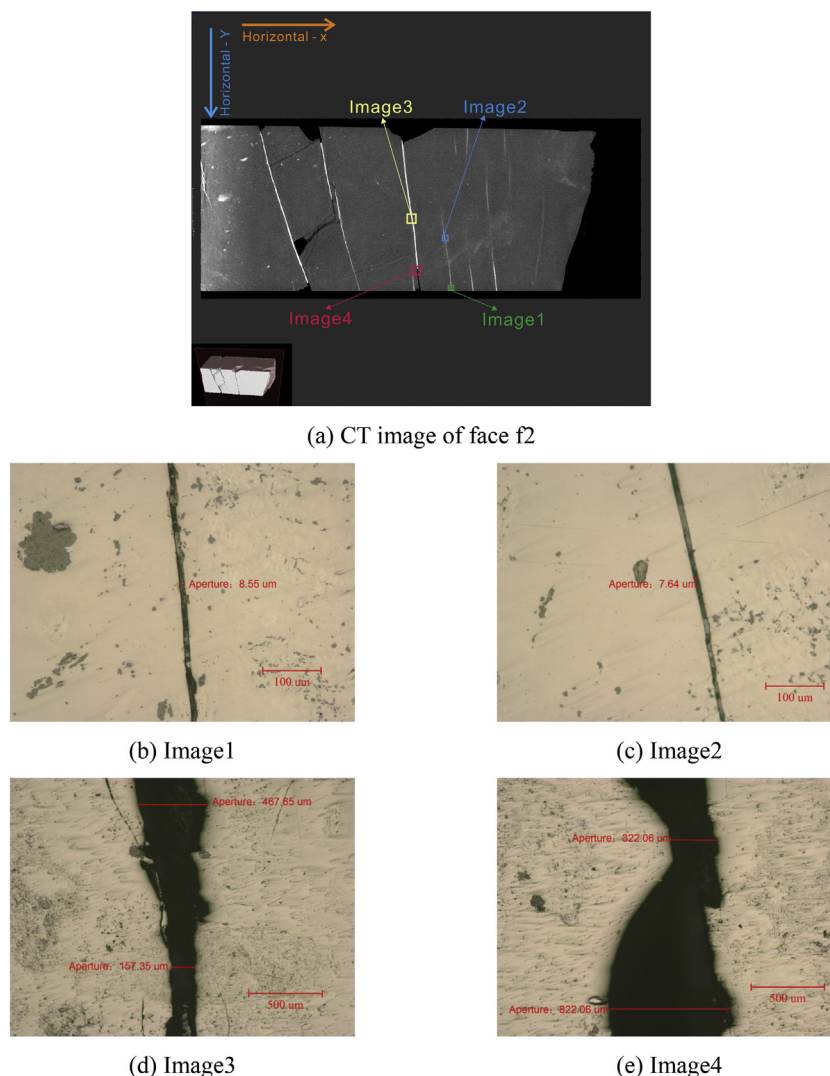


Fig. 4. Fracture distribution and fracture aperture distribution on the surface of coal sample face f2 as recovered through CT imaging (a) and optical microscopy (b-e).

Liu, 2019; Zhao et al., 2019), and also affects the matrix permeability. MICP was conducted with a Mercury Injection Porosimeter (instrument: Micromeritics AutoPore IV9510) to investigate the pore size distribution of the coal sample, performed on cubic samples $\sim 1 \times 1 \times 1 \text{ cm}^3$ after drying at 60°C for 48 h to remove moisture (Yang et al., 2017). The surface tension of the mercury is 485 mN/m with a contact angle of 130° (Gan et al., 1972; Sun et al., 2017). The corresponding pore-throat distribution was calculated using the Washburn equation (Washburn, 1921).

2.2. Experimental set-up

The experimental set-up for this study was designed to enable the measurement of strain induced by the mechanical compression and shrinkage or swelling resulting from the gas adsorption (Liu and Harpalani, 2013). A schematic of the experimental set-up is illustrated in Fig. 2. The set-up comprises a gas pressure controller, pressure monitoring system, high-pressure vessel and a data acquisition system to monitor the strains. The vessel was placed in a thermostatic chamber to eliminate the thermal impact on the deformation of the coal sample.

2.3. Experimental procedure

Strain gauges were applied to the sample to measure the spatial distribution of coal deformation history, and to resolve the coupling between shrinkage deformation and dynamic gas injection (Kiyama et al., 2011; Liu et al., 2016b). The experimental set-up using strain gauges was designed and employed to measure the sorption-induced strain in the horizontal direction (Liu and Harpalani, 2013; Sang et al., 2017). Four strain gauges were installed on the designated surface of the sample so as to monitor the local linear strains in the horizontal direction (Fig. 3a). According to the distribution of natural fractures, Gauge A was used to measure the average deformation of the whole sample; Gauge B was adopted to measure the average deformation of the part of the sample that crosses a single dominant (large) fracture; Gauge C was arranged to measure the average matrix deformation of one part of the sample that contains no fractures, and; Gauge D was used to measure the average deformation of one part of the sample that crosses a small fracture, as shown in Fig. 3 (a). After the attachment of gauges, the instrumented surface (f2) was covered in soft silicone rubber adhesive to protect the strain gauges for the long-time deformation test and to prevent gas from entering the sample across this surface, as shown in Fig. 3 (b). The entire assembly was then placed in the pressure vessel, as illustrated in Fig. 2. After thermally

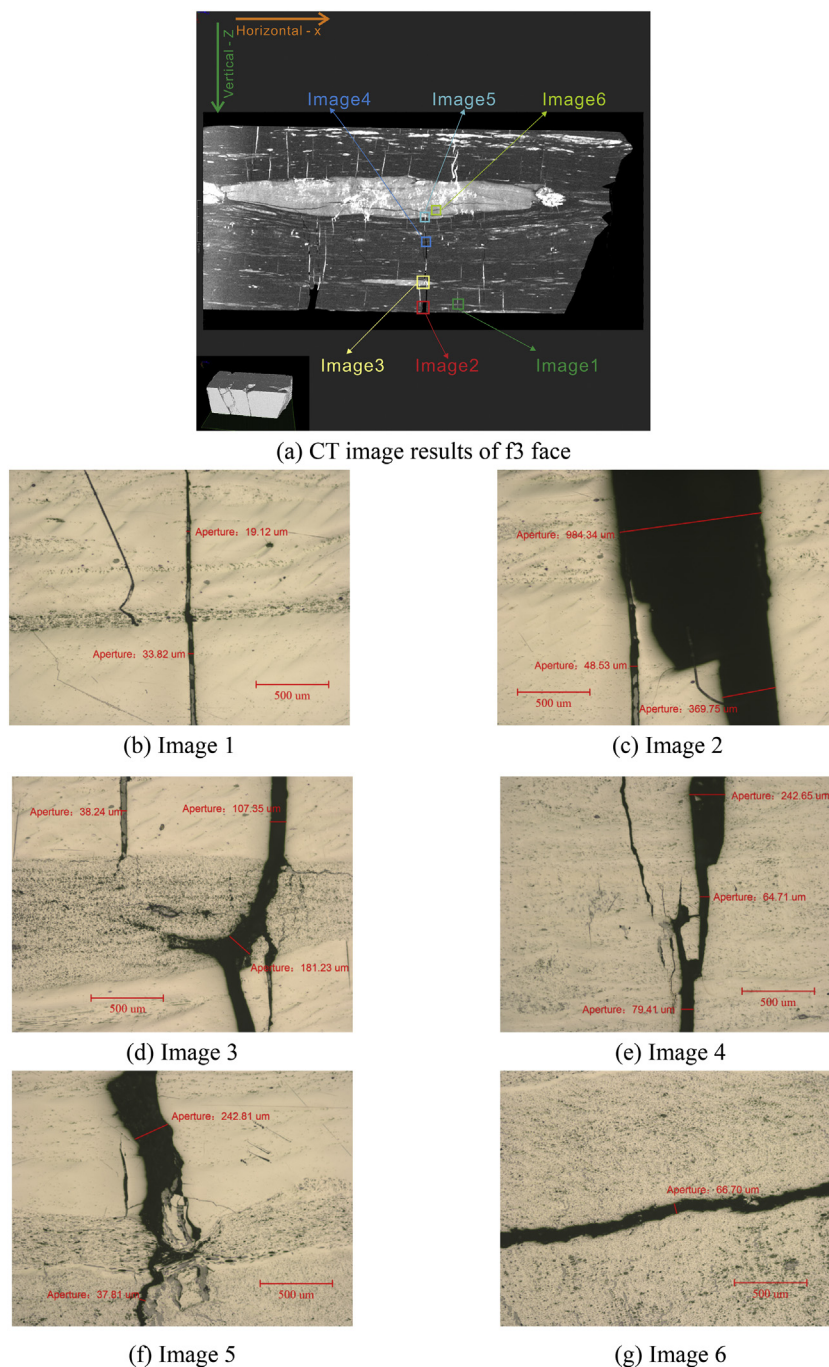


Fig. 5. Fracture distribution on the surface of face f3 of the coal sample as recovered through x-ray CT imaging (a) and optical microscopy (b-g).

equilibrating, pore pressure was gradually increased from 0 to 1.0 MPa, at a rate of 0.1 MPa/10s, and then remained constant at 1.0 MPa. The temperature was maintained at 26.5 ± 0.3 °C during the entire experiment.

3. Experimental results and discussion

The characterized samples are subject to dynamic experiments involving sorption and measuring the active generation of surface strains at a spectrum of relevant scales. The permeability of the matrix system is estimated to be in the nano-Darcy range, according to the MICP tests. These define the characteristic of the matrix system and the fracture system for later analysis. The distributed strain gauges directly observed the evolution of average matrix strain for the full extent of the

sample that is covered by the strain gauge.

3.1. Description of natural fracture systems

Coal is conceptualized as a dual porosity/permeability system containing porous matrix surrounded by fractures. In this study, the cleat system, fractures, joints, and faults are all included in the fracture system (Shi et al., 2018). It is commonly assumed that Darcy flow is a result of flow in the fracture system and that the contribution of Darcy flow in the coal matrix can be neglected (Purl et al., 1991). Thus the bulk permeability of a coal sample is a function of the fracture system (Palmer, 2009; Pan and Connell, 2012; Ried et al., 1992). Therefore, the distribution of the fracture network determines the primary pathway to gas flow, but the connectivity of this fracture system is influenced by

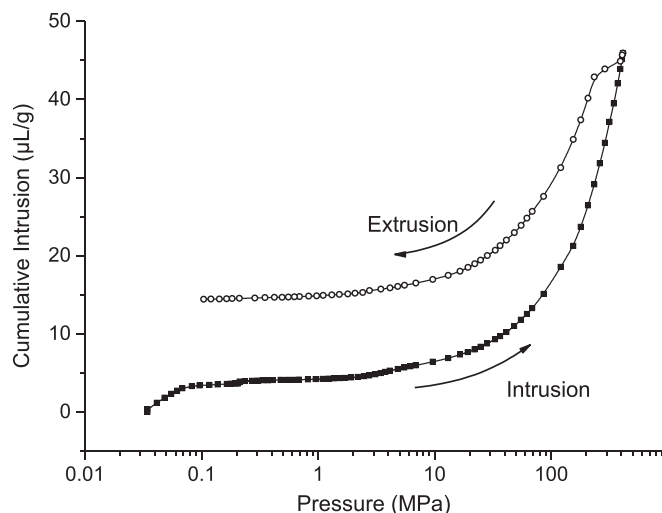


Fig. 6. Mercury intrusion and extrusion volumes with pressure for the coal sample.

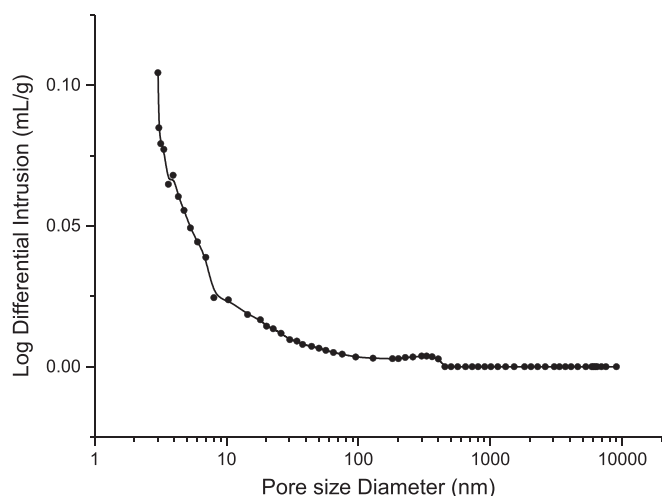


Fig. 7. Log differential intrusion volume vs. pore size diameter for the coal sample.

deformation of the matrix.

The distribution of fractures on face f2 is mostly in the Horizontal ($-Y$) direction, with most of these fractures filled with calcite (white lines), as shown in Fig. 4 (a). We selected two representative fractures for analysis – a small one and a large one. The small fractures show a largely uniform aperture which is of the order of $\sim 8 \mu\text{m}$ within the field of view, as shown in image1 and image2 of Fig. 4 (b) and (c). Conversely, the single large fracture exhibits a relatively tortuous topology with a non-uniform aperture ranging from ~ 150 to $\sim 850 \mu\text{m}$ within the field of view (see image3 and image4 in Fig. 4 (d) and (e)).

The distribution of fractures on face f3 is mostly in the Vertical-Z direction, with most of these fractures filled with calcite, as shown in Fig. 5 (a). The distribution of fractures on face f3 is more complex than that on face f2, as face f2 is perpendicular to bedding. Three representative fractures were selected for analysis. These are: (i) a small fracture as shown in image1 in Fig. 5 (b) which flanks the small fracture in Fig. 4 (b) and (c), (ii) a large fracture as shown in image2–5 in Fig. 5 (c–f) which flanks the large fracture in Fig. 4 (d) and (e), and (iii) a fracture in a rock parting, as shown in image 6 in Fig. 5 (g).

3.2. Pore size distribution of matrix system from MICP tests

MICP is an effective tool in providing quantitative evaluation of pore structure. Fig. 6 shows the cumulative intrusion and extrusion volume versus the applied pressure for the studied coal sample. The pore size distribution is presented in Fig. 7. A significant volume of mercury intrudes the connected pores in the coal sample with an increase in applied pressure, suggesting that the sizes of some pores are in the meso-macropore range.

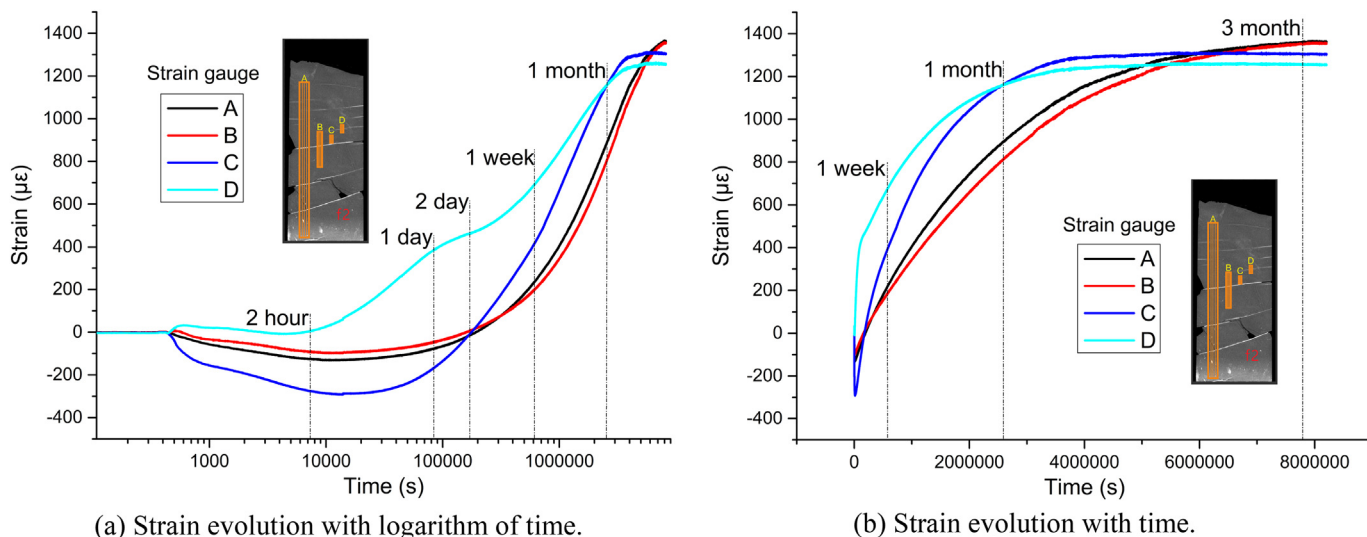
The average pore throat size, indexed by volume (50%) of the coal sample, is less than 10 nm. The permeabilities of the connected pore network for pore throat diameters (matrix system) in the range 4–10 nm is in nano-Darcy range (Hu et al., 2015; Katz and Thompson, 1986; Sun et al., 2017; Yang et al., 2017).

3.3. Changes of coal local linear strain with time

The fracture voids are devoid of filling and reject the attachment of the strain gauges – thus direct compression or extension cannot be measured on this section. The strain data acquired from the measurement is only able to evaluate the evolution of *average* matrix strain for the full extent of the sample that is covered by the strain gauge. Fig. 8 presents the evolution of the coal strains with time during injection of CH_4 . Strain gauges A, B, C and D were located respectively on different parts of the sample face to sample the response of the different fracture distributions. Strain gauge A recorded the deformation of the overall sample; strain gauge B recorded the deformation of the part of the sample that crosses a large fracture; strain gauge C recorded the deformation of a small part of the sample that contains no fractures; and strain gauge D recorded the deformation of a small part of the sample that crosses a small fracture. The evolution of strains A and C with time can be divided into three stages, while the evolution of strains B and D with time can be classified into four stages. The monitoring results exhibit nearly the same trends. First, local linear strain increased rapidly from the initiation of gas injection (except for strain gauges A and C). Following that, the local linear strains begin to decrease gradually to reach a minimum, but the rate and magnitude of this decay was inconsistent. Subsequently, local linear strains begin to gradually increase and finally reach a plateau, with this process lasting for the extraordinarily extended period of two to three months. During this entire process, the maximum shrinkage strain was of the order of $-292\mu\epsilon$, and the maximum swelling strain of about $1365\mu\epsilon$.

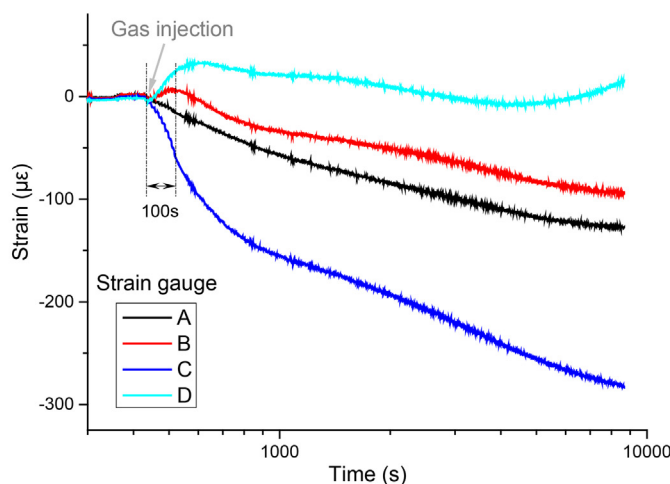
The deformation of different part of the sample with different fracture distributions show different evolutionary trends. Strain gauges A and B (black line and red line in Fig. 8), that cross large fractures and sample a relatively large tributary area of solid matrix, show nearly the same evolutionary trend. They reach nearly the same magnitude (about $-100\mu\epsilon$) of maximum shrinkage strain and at the same time. Then these two local linear strains gradually increased and finally reached a plateau with a maximum swelling strain of $\sim 1360\mu\epsilon$ – but taking nearly three months to reach this equilibrium state. While for strain gauge D (cyan line in Fig. 8), which measures the deformation of a small part of the sample that crosses a small fracture, the evolutionary trend is totally different from that of strain gauges A and B. The maximum shrinkage strain is only approximately $-9\mu\epsilon$, and it takes only slightly more than one month to reach the final plateau. For strain gauge C (blue line in Fig. 8), which measures the deformation of only a small part of the sample and that contains no fractures, the time taken to reach the final plateau is nearly the same as that for strain gauge D, but it shows a clear shrinkage displacement with a maximum shrinkage strain of approximately $-292\mu\epsilon$.

These results show the different average matrix strain evolution rates and magnitudes for different parts of the sample. For the small strain gauges (C and D) close to the fracture, the equilibrium time is shorter than that for the larger gauges (A and B) measuring the response of a larger tributary area of the matrix that has no obvious fractures



(a) Strain evolution with logarithm of time.

(b) Strain evolution with time.



(c) Strain evolution –with logarithm of time, in the first 2 hours.

Fig. 8. Evolution of strain with time during injection of CH₄.

present. It takes nearly two to three months to reach the final equilibrium state of this system, which is much longer than the typical “equilibrium time” used in experimental measurements of permeability. However, most significantly, the average matrix strains do not evolve monotonically throughout the experiment, which is generally assumed in most permeability models that do not accommodate this spatial evolution in deformation and strain – this is a fundamental observation that is important in correctly characterizing such systems. Thus, to further analyze these interaction mechanisms, we constructed a numerical model to analyze the experimental results.

4. Numerical model implementation and validation

In the following, a fully coupled numerical model is implemented to understand principal active mechanisms of sorption-induced straining and to reconcile laboratory observed response. Specifically, we match modeling results with laboratory measurements, to numerically simulate the spatial distribution of strain, to explore the dynamics of stress transfer between the matrix and the fracture, and to model the evolution of coal permeability from initial to ultimate equilibrium.

4.1. Governing equations of coal deformation

We assume that both the matrix and the fracture are homogeneous

and elastic, and that deformation obeys Hooke's law. The deformation of a homogeneous, isotropic, and uniform elastic medium can be described by the Navier-type equation both in the matrix region and the fracture region (Wu et al., 2010):

$$Gu_{i,kk} + \frac{G}{1 - 2\nu}u_{k,i,k} - \alpha p - K\varepsilon_L \frac{p_L}{(p + p_L)^2}p_i + f_i = 0 \quad (1)$$

where $G = E/2(1 + \nu)$, $K = E/3(1 - 2\nu)$, G is the shear modulus, u_i is the component of displacement in the i -direction, ν is the Poisson's ratio, α is the Biot's coefficient, K is the bulk modulus, E is the Young's modulus, ε_L is the Langmuir volumetric strain constant representing the volumetric strain at infinite pore pressure only in the matrix region, p_L is the Langmuir pressure at which the measured volumetric strain is equal to $0.5 \varepsilon_L$ only in the matrix region, p is the pore pressure, and f_i is the component of body force in the i -direction. Specifically, these parameters represent different properties (e.g., G_m is the shear modulus of the matrix, G_f is the shear modulus of the fracture) for the matrix and fracture.

4.2. Mass balance equation

The flow in both fracture and matrix is assumed to be governed by Darcy's law. Although it is commonly assumed that Darcy flow predominates in the fracture system, and diffusion may dominate in the

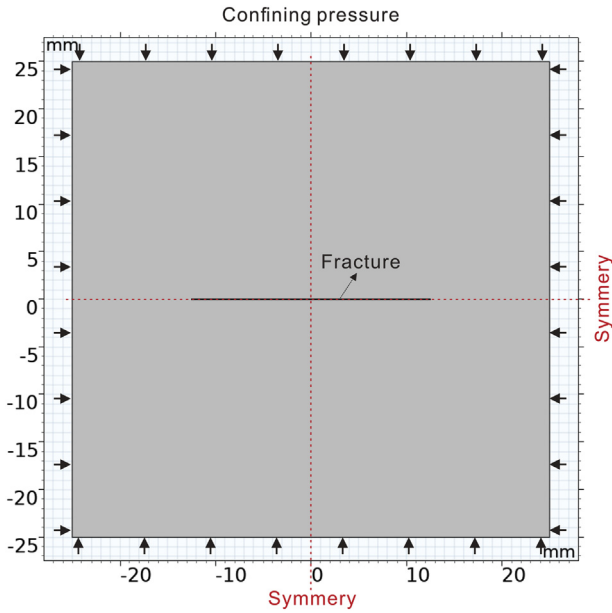


Fig. 9. Geometry of the 2D symmetric model together with boundary conditions.

coal matrix (Purl et al., 1991) – because there are equivalencies between permeability and diffusivity.

(1) Governing equation for gas flow within matrix.

According to our previous work (Wei et al., 2019a), for the case of an ideal absorbing gas, the governing equations for matrix and fracture are:

$$\frac{d}{dt} \left(\phi_m p_m \frac{M}{RT} + (1 - \phi_m) \rho_s \rho_a \frac{V_L p_m}{p_m + P_L} \right) - \nabla \cdot \left(\rho_g \frac{k_m}{\mu} \nabla p_m \right) = Q_s \quad (2)$$

where ϕ_m is the porosity of matrix system, p_m is the gas pressure in the matrix, M is the molecular mass of gas, R is the universal gas constant, T is the absolute gas temperature, ρ_s is coal density, ρ_a is gas density at atmospheric pressure, V_L is the Langmuir volume constant, ρ_g is the gas density, μ is the dynamic viscosity of the gas, k_m is the permeability of matrix system, and Q_s is the gas source or sink.

(2) Governing equation for gas flow within fracture.

Gas transfer through the fractures is governed by the basic mass conservation relation of the fracture system, but is rarely used in

models of matrix-fracture interaction, due to the rapid equilibration within the fracture (Liu et al., 2018). Thus, the pore pressure in the fracture p_f is assumed to be increased gradually within 100 s from the application of the initial pore pressure p_{m0} (0.1 MPa) to a constant p_1 (1.0 MPa).

$$p_f = \begin{cases} p_1 & t \geq 100s \\ t * 0.009 * p_1 + p_{m0} & t < 100s \end{cases} \quad (3)$$

The fracture permeability is defined by the well-known “cubic law” (Witherspoon et al., 1980) and the fracture permeability ratio can be expressed as:

$$\frac{k_f}{k_{f0}} = \left(1 + \frac{\Delta b}{b_0} \right)^3 \quad (4)$$

where k_f is the fracture permeability, k_{f0} is the initial fracture permeability, b_0 is initial fracture aperture and Δb is the change in fracture aperture. The field equations are implemented and solved using COMSOL Multiphysics.

4.3. Model implementation and validation

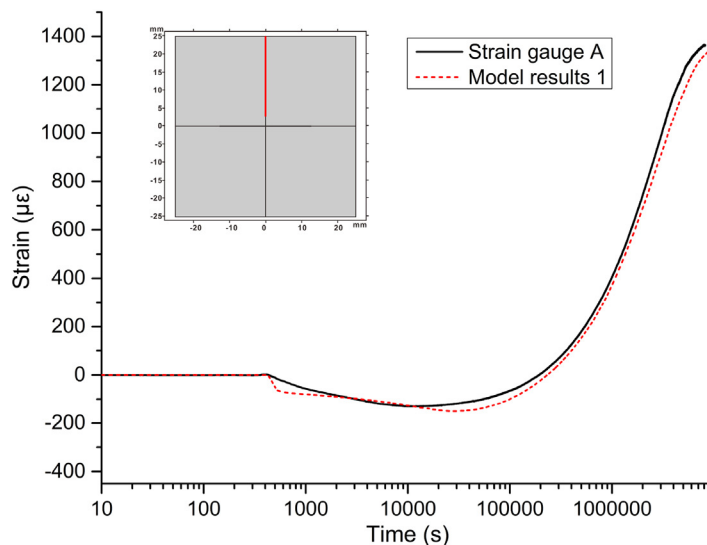
As shown in Fig. 9, the sample is simplified as a 2D model with both length and height of 50 mm. A fracture is distributed in the middle of the sample (black line in Fig. 9), and the length of the fracture is 25 mm with an aperture of 45 μm . Injection of CH_4 is taken as an example to illustrate how mass and stress transfer between the matrix and the fracture are coupled under the constant confining pressure condition and the distribution of the heterogeneous strains inside the coal sample under the influence of gas-sorption determined. It is assumed that the gas is injected into the center of the model. No flow conditions are applied to all external boundaries with mass exchange between the matrix and fracture occurring on the connecting boundary. The interface boundary condition between the matrix and fracture domains is one of pressure continuity across interior boundary (Wei et al., 2019a). The mass transfer between the matrix and fracture is applied at the interface between these two systems. The applied confining pressure and pore pressure increase gradually within 100 s to a constant (1.0 MPa). Relevant initial values of parameters for fracture and matrix system in the simulation model are listed in Table 1. Simulated strain evolution in different part of the sample is compared with the experimental results at the same stress conditions in Fig. 10. Results show that the modeled evolution of strain is in good agreement with the strains that are directly measured by strain gauges A and C on the sample.

4.4. Analysis of pore pressure, stress and strain evolution

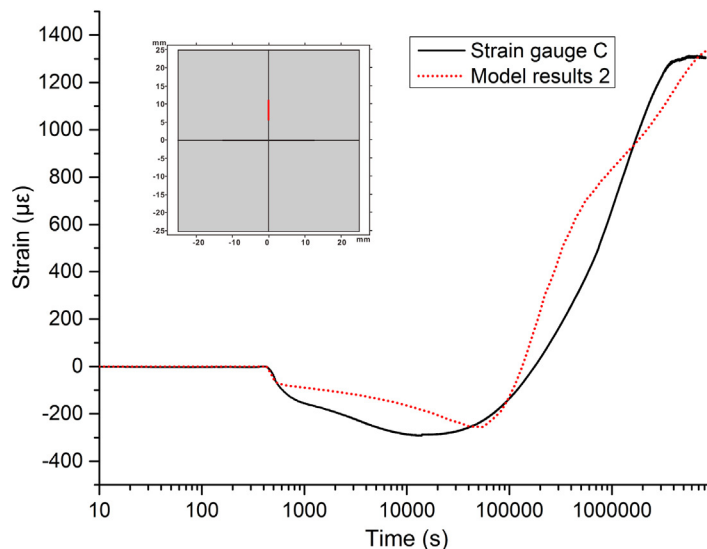
Fig. 11 shows the distribution of pore pressure, stress and strain in the matrix at different times during CH_4 injection. The correlation

Table 1
Property parameters of explicit model.

| Parameters | Value | Source |
|--|----------------------------------|---|
| Young's modulus of matrix, E_m | 5.5 GPa | Modified according to Liu et al., 2018 |
| Young's modulus of fracture, E_f | 0.05 GPa | Modified according to Liu et al., 2018 |
| Langmuir volume of matrix, V_L | 0.017 m^3/kg | Modified according to Liu et al., 2018 |
| Langmuir pressure of matrix, p_L | 3 MPa | Modified according to Liu et al., 2018 |
| Langmuir volumetric strain, ϵ_L | 0.0093 | Zhang et al., 2018 |
| Poisson's ratio of matrix, ν_m | 0.35 | Zhang et al., 2018 |
| Poisson's ratio of fracture, ν_f | 0.35 | Zhang et al., 2018 |
| Biot's coefficient, α | 0.7 | Zhang et al., 2018 |
| Coal density, ρ_s | 1500 kg/m^3 | Zhang et al., 2018 |
| Viscosity of the gas, μ | 1.84×10^{-5} | Constant |
| Porosity of matrix | 0.02 | Zhang et al., 2018 |
| Permeability of matrix, k_m | $2.3 \times 10^{-21} \text{m}^2$ | Modified according to Wei et al., 2019b |
| Initial fracture aperture, b_0 | 45 μm | Experimental test |
| Injection pressure, p_1 | 1 MPa | Experimental setting |



(a) Comparison between model and experimental results for strain gauge A.



(b) Comparison between model and experimental results for strain gauge C.

Fig. 10. Comparison between the model results and experimental data.

between the evolution of strain (solid line) and pressure (dotted line) at the red line in the sample (model results 1 in Fig. 10 (a)) at different time is presented in Fig. 12.

Prior to injection, the sample is at equilibrium (pressure, stress and mass contents) and no disequilibrium interaction between the matrix and the fracture occurs (Fig. 11 (1) and Fig. 12 (initial state)). Post-injection, a series of processes initiate. Confining pressure and pore pressure are increased gradually from 0 to 1.0 MPa within 100 s, and then maintain constant at 1.0 MPa. First, gas rapidly invades the fracture due to its relatively high permeability, then the distribution of the gas pressure in the fracture remains constant, leading to a pressure difference between the matrix and the fracture (Fig. 11 (2) and Fig. 12). Next, with the gas contacting the fracture surface and diffusing into the matrix, local strains evolve in the matrix due to both the gas adsorption and the increasing gas pressure, which occurs only locally in the direct vicinity of the fracture. These coupled diffusion/sorption phenomena

determine the localized deformation within the ensemble matrix-fracture system. Under this condition, the matrix near the fracture swells, causing the fracture to narrow/compact, and shrinkage of the matrix occurs far from the fracture due to the redistribution of stress (Fig. 11 (2) and Fig. 12 (2 h, 1 day)). As gas injection continues, the swelling zone widens as the gas diffuses further into the matrix of the sample (Fig. 11 (2–5)). The gas pressure front propagates throughout the matrix until a new equilibrium state between the fracture and the matrix is attained (Fig. 11 (6) and Fig. 12 (final equilibrium state)). The results show that during this non-equilibrium condition, the swelling of the matrix near the fracture will cause not only compaction and narrowing of the fracture, but also shrinkage of the matrix that is distant from the fracture. The above analysis demonstrates that many permeability measurements in ultra-low permeability rocks may have been conducted under the non-equilibrium conditions, suggesting that the interactions between the matrix and the fracture must be taken into

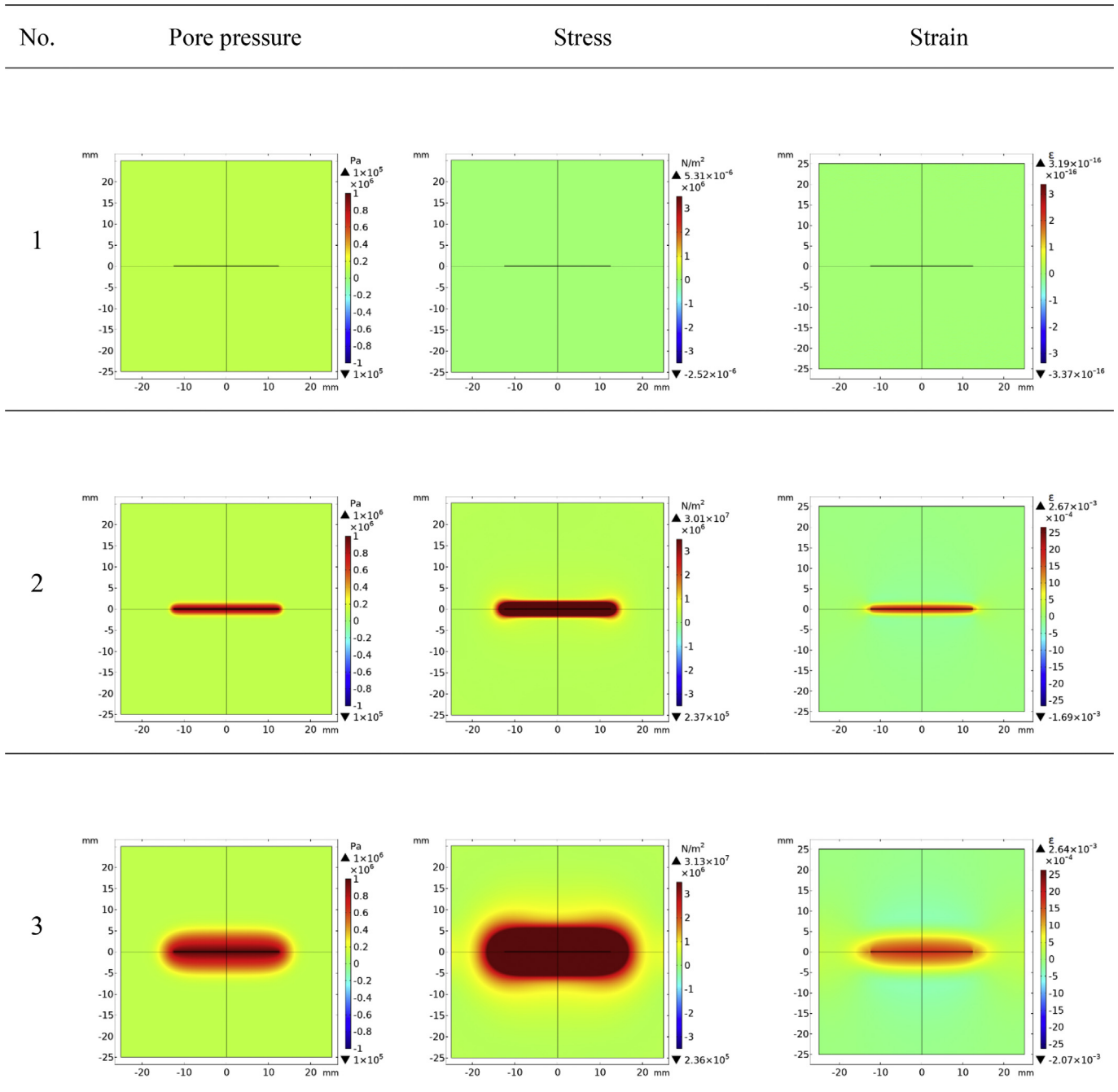


Fig. 11. Distribution of pore pressure, stress and strain in matrix at different times during CH₄ injection.

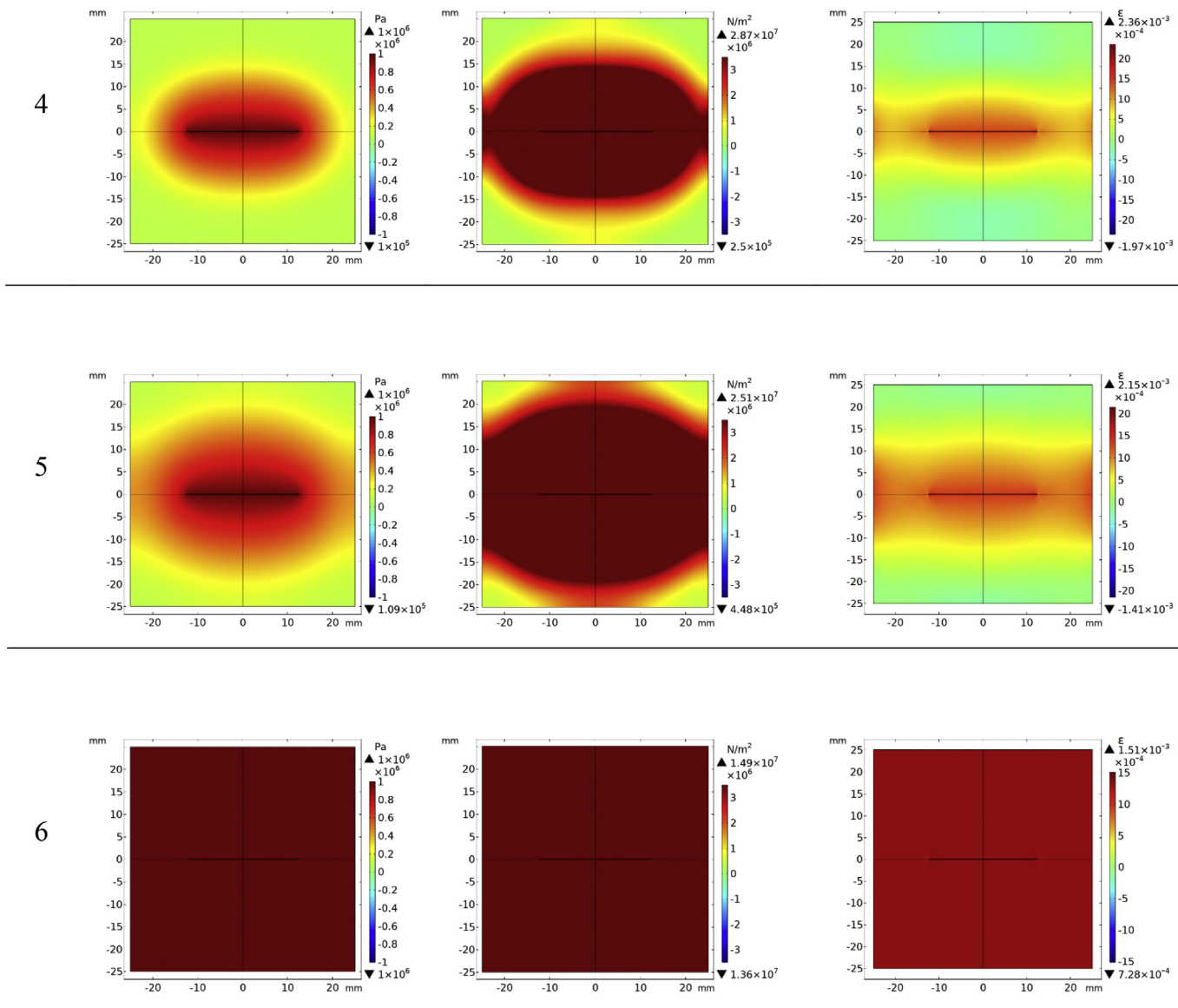
consideration to fully understand the stress transfers and to correctly interpret the permeability evolution. This indicates a significant gap between most of the existing equilibrium-assumption permeability models and laboratory experiments.

4.5. Analysis of permeability evolution

In the prior experimental and numerical observations of ultra-low permeability coal, strains are demonstrated to change throughout the entire period of experiment and in turn alter the fracture aperture-determining the evolution of non-equilibrium fracture permeability.

Fig. 13 shows the evolution of permeability based on our model. The behavior can be divided into three stages under our particular experimental conditions: (1) Stage 1 - permeability decrease. During the gas injection period, the applied confining pressure and pore pressure

increase gradually from 0 to 1.0 MPa within 100 s and then remain constant at 1.0 MPa. There is an aperture reduction induced by the influence of Biot's coefficient and a time delay for the equilibration of pore pressure both in the fracture system and the matrix system. As the gas diffuses from fracture to matrix, the gas pressure in the matrix slowly increases. This process leads to the swelling of the matrix adjacent to the fracture (Fig. 11 (2–3) and Fig. 12 (2 h, 1 day)). In addition, this localization of swelling reduces the fracture aperture and further causes a decrease of permeability; (2) Stage 2 - permeability recovery. As the gas front propagates from the vicinity of the fracture wall to the exterior region (distant from the fracture), the matrix strain becomes more uniform (Fig. 11 (4–5) and Fig. 12 (1 week, 2 weeks)). At this stage, the effect of local strain in the matrix vanishes with the permeability beginning to gradually increase and finally reaching a plateau, which lasts for the extraordinarily extended period of three



(1) Initial state - before gas injection; (2) 2 hour - after gas injection; (3) 1 day - after gas injection; (4) 1 week - after gas injection; (5) 2 week - after gas injection; (6) Final equilibrium state.

Fig. 11. (continued)

months; (3) Stage 3 - permeability stabilization. When a final equilibrium state is reached, the effect of local strain completely vanishes (Fig. 11 (6) and Fig. 12 (final equilibrium state)) and the permeability remains unchanged. Thus, permeability changes during non-equilibrium states are the consequences of complex interactions between coal matrix and fracture. Similar process will also prevail during gas extraction (e.g., coalbed methane production, shale gas extraction) rather than injection, which is studied here.

5. Conclusions

In this study, we directly measured and analyzed the strain evolution of a prismatic coal sample during gas injection to experimentally determine how coal strain/permeability evolves from initial to ultimate equilibrium. The following conclusions can be drawn:

- The experimental results show that the interval from initial to final equilibrium is much longer than that generally assumed and reported in the literature. The final equilibrium state (pressure, stress and mass contents) for the matrix system may take an unusually long time and therefore the current permeability measurements may reflect non-equilibrium permeability evolutions. Although the disequilibrium sorbing masses are potentially small, the innate sensitivity of permeability in the matrix-fracture system makes permeability estimates potentially quite sensitive to this disequilibrium state.
- During the non-equilibrium condition, the distribution of deformations in different parts of the sample is not uniformly distributed in space or in time – rather they are strongly related to the fracture geometry and related gas saturation conditions. The swelling of the matrix near the fracture will cause not only compaction and narrowing of the fracture, but also shrinkage of the matrix that is

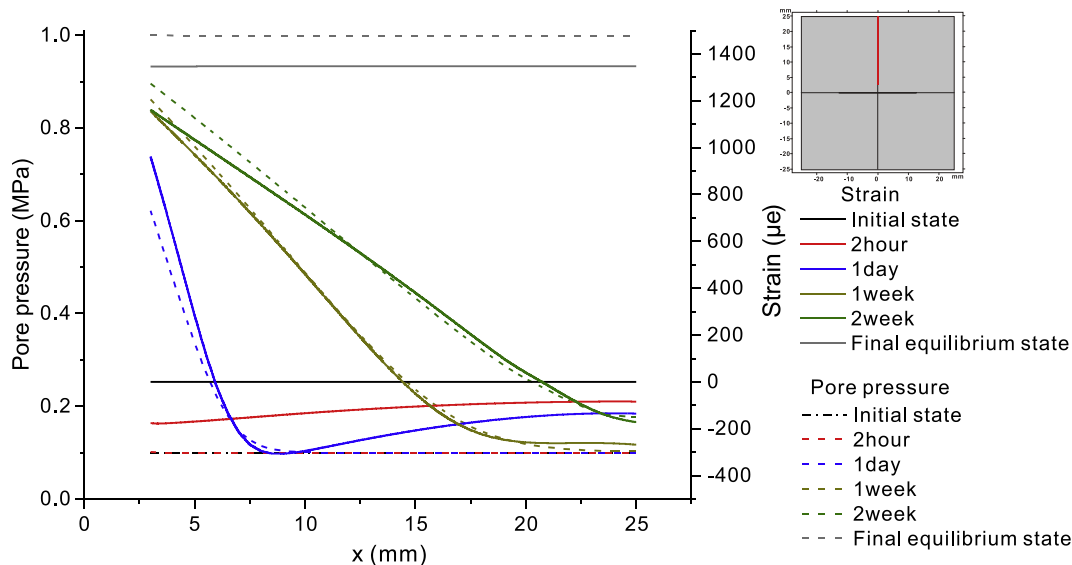


Fig. 12. Correlation between gas pressure and strain evolution.

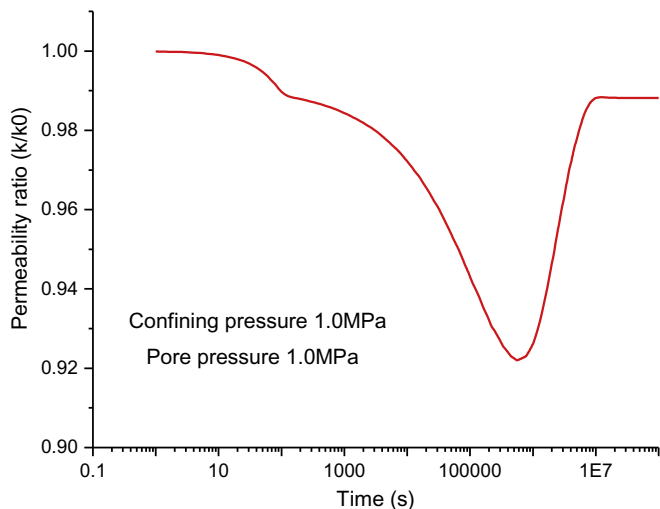


Fig. 13. Evolution of fracture permeability ratio.

distant from the fracture under constant confining pressure condition.

- Coal strains change through the entire duration of the experiment and in turn alter the fracture aperture, with the evolution of non-equilibrium strain/permeability determined by the matrix-fracture interactions while that of equilibrium permeability is defined by the overall behavior.
- There is a gap between most of the existing equilibrium-assumption permeability models and laboratory experiments. This gap is determined mainly by a lack of recognition of this prominent non-equilibrium state and the impact of matrix-fracture interactions including sorption-induced swelling/shrinking for all gas pressures. Permeability changes through this non-equilibrium state are the consequence of complex interactions between coal matrix and fracture and may be much more common and of much longer duration than previously suspected.

Declaration of Competing Interest

The authors declare that they have no conflict of interest.

Acknowledgments

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