



**Changes of Sea Level
and Poleward Ocean Heat Transport as
Potential Causes for the Late Ordovician
Glaciation**

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Objectives

- Using an Atmospheric General Circulation Model to
 - Evaluate the sensitivity of the Late Ordovician climate to different paleogeographies, atmospheric $p\text{CO}_2$ values, sea levels, and poleward ocean heat transport.
- Using an Ice Sheet Model to
 - Investigate the threshold that led to the formation of ice sheets.
- Using a Ocean General Circulation Model to
 - Evaluate the impact of paleogeographic changes, atmospheric $p\text{CO}_2$ values, and sea level changes on global poleward ocean heat transport and the initiation of ice sheets during the Late Ordovician.

Methods - 1

- General Atmospheric Circulation Model – GENESIS v.2.0
- T31 spectral resolution (~400 km) for atmosphere
- 2° x 2° surface resolution
- 50-m slab ocean, dynamic sea ice model
- 40-year runs

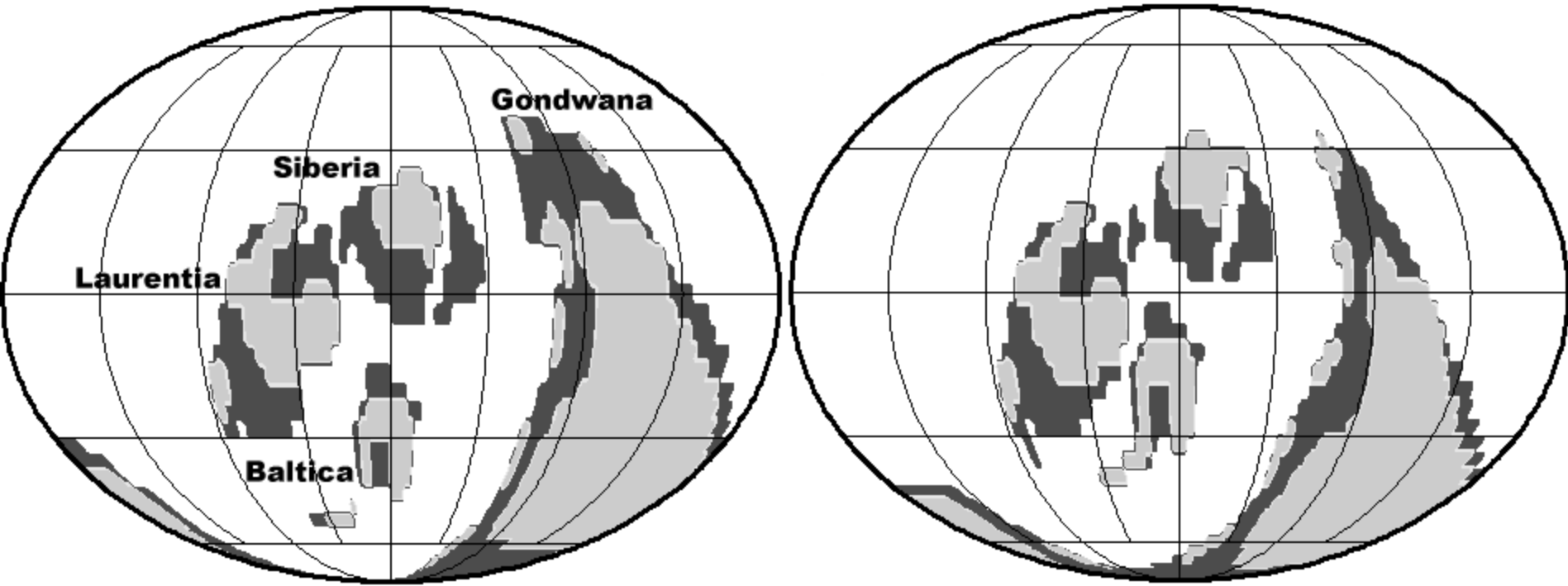
Boundary Conditions

| Boundary Conditions | How varied |
|------------------------------|---|
| Land-Sea distribution | Paleogeography of Scotese and McKerrow (1991); shoreline position after Scotese (1997) |
| Topography | 250 m for coastal grid points and 500m for all other land grid points |
| Orbital parameters | “Cold-Summer Orbit” eccentricity: 0.06 obliquity: 22. precession: 270. perihelion to N.H. vernal equinox |
| Solar luminosity | 4.5% reduction of present day |
| Vegetation and soil type | No vegetation with intermediate soil color values |
| Atmospheric pCO ₂ | 5040 ppm (18x PAL); 4200 ppm (15x PAL); 3360 ppm (12x PAL); 2800 ppm (10x PAL); 2240 ppm (8x PAL) |

Paleogeographic Changes

Caradocian (454 Ma)

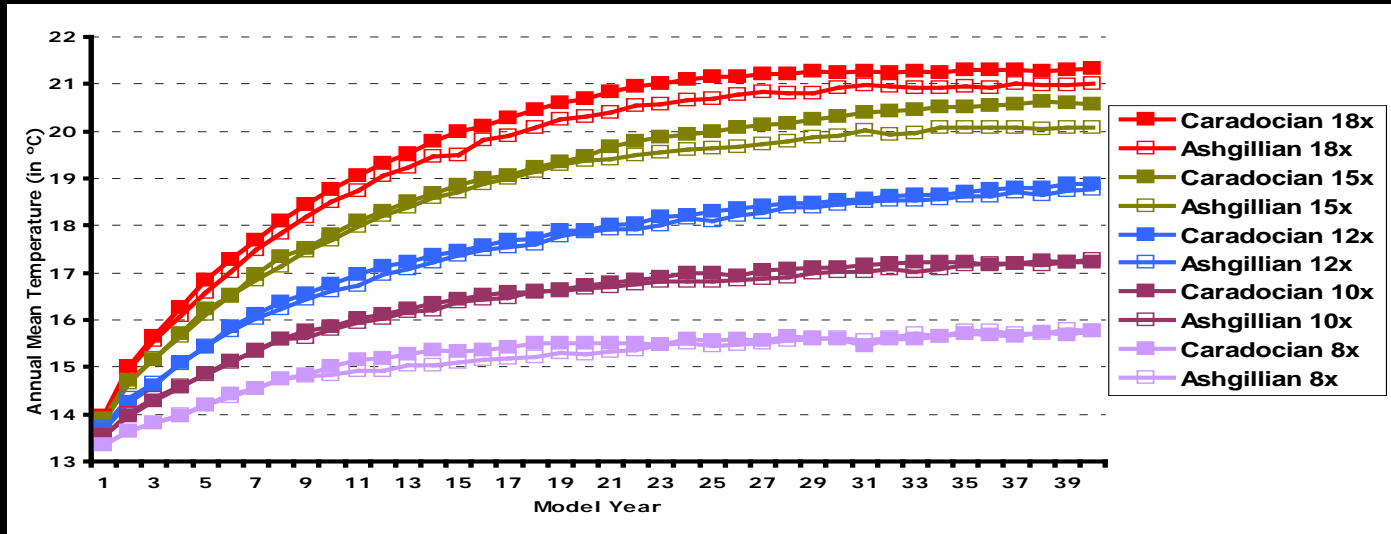
Ashgillian (446 Ma)



(after Scotese and McKerrow 1990, 1991; Scotese 1997)

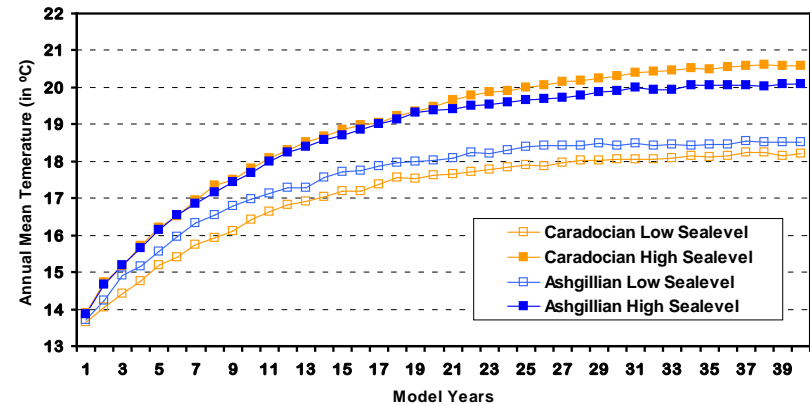
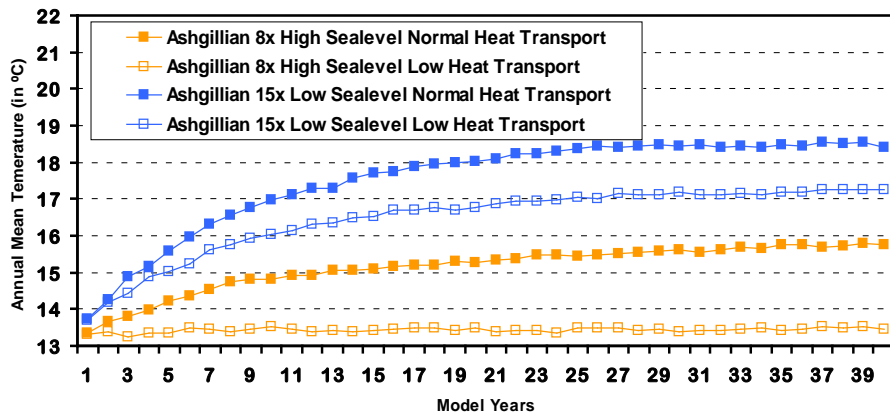
Annual Mean Temperatures

Paleogeography and pCO₂



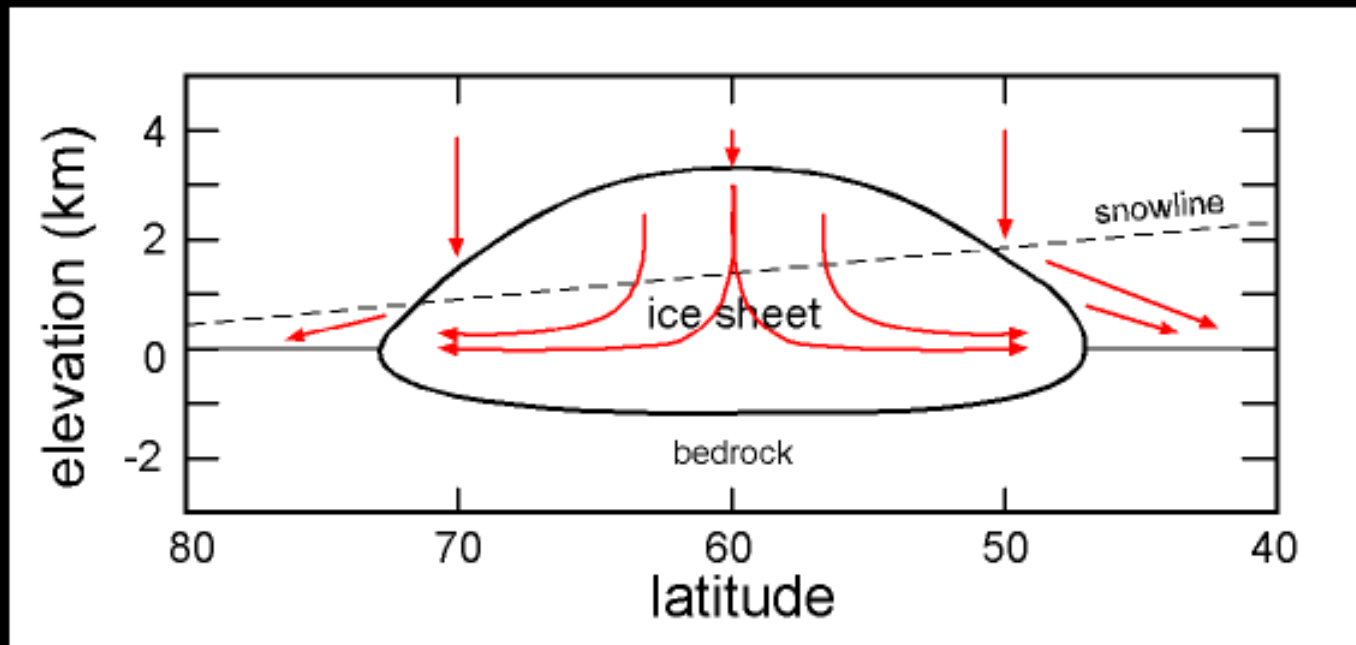
Poleward heat transport

Sea level



Methods 2 - Ice Sheet Model

- 3-dimensional model
- $1^\circ \times 1^\circ$ longitude-latitude grid
- Surface mass balance: Degree-day
- 2-km bedrock with vertical heat diffusion



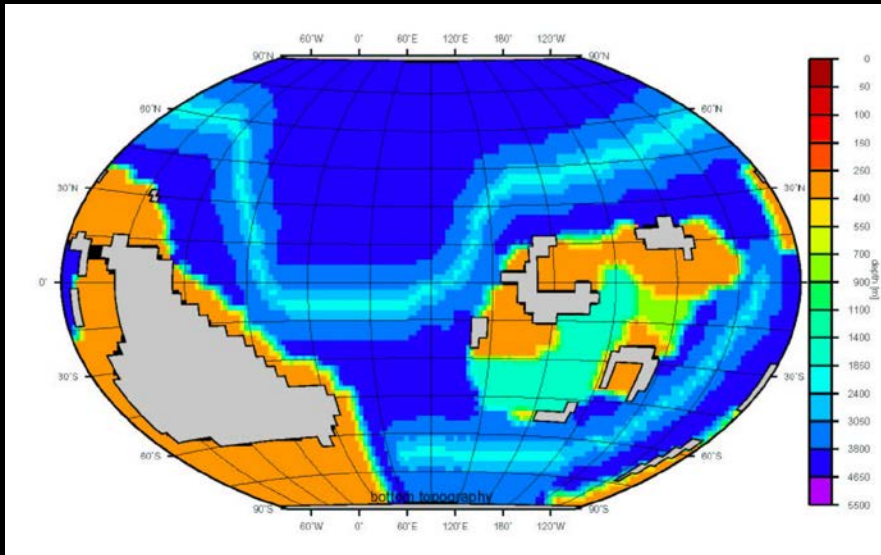
Ice Sheet Model

| Simulation | Caradocian | Ashgillian |
|--|-------------------|-------------------|
| 18x High Sealevel Normal Heat Transport | No glaciation | No glaciation |
| 15x High Sealevel Normal Heat Transport | No glaciation | No glaciation |
| 15x Low Sealevel Normal Heat Transport | No glaciation | No glaciation |
| 15x Low Sealevel Low Heat Transport | Glaciation | No glaciation |
| 12x High Sealevel Normal Heat Transport | No glaciation | No glaciation |
| 10x High Sealevel Normal Heat Transport | No glaciation | No glaciation |
| 8x High Sealevel Normal Heat Transport | No glaciation | No glaciation |
| 8x Low Sealevel Normal Heat Transport | Glaciation | Glaciation |
| 8x High Sealevel Low Heat Transport | Ice Sheets | Glaciation |

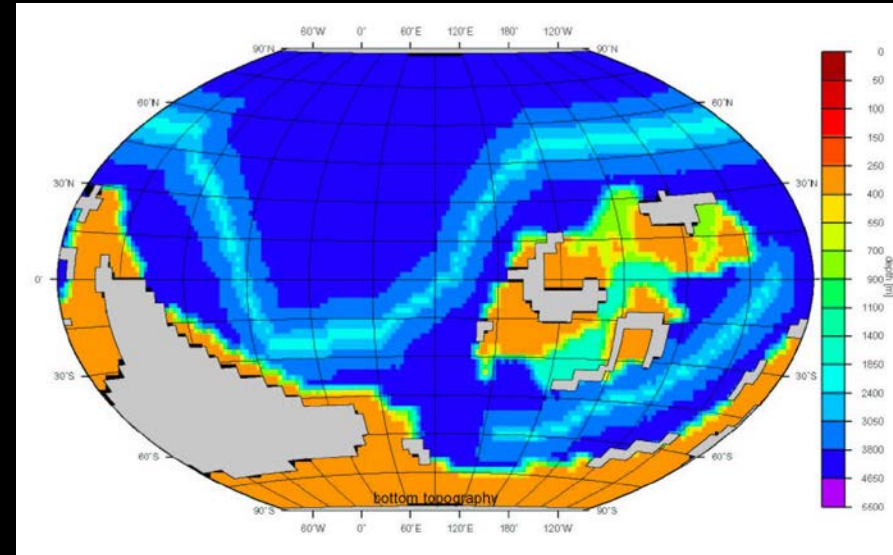
Methods 3 – Ocean Circulation Model

- Modular Ocean Model v.2.2
- 4° x 4° longitude-latitude grid and 16 unequally spaced vertical layers
- Atmospheric forcing from GENESIS simulations
- All runs are 2000 model years long, with five-fold acceleration in the deep layers (i.e., deep ocean is effectively run for 10000 yr).

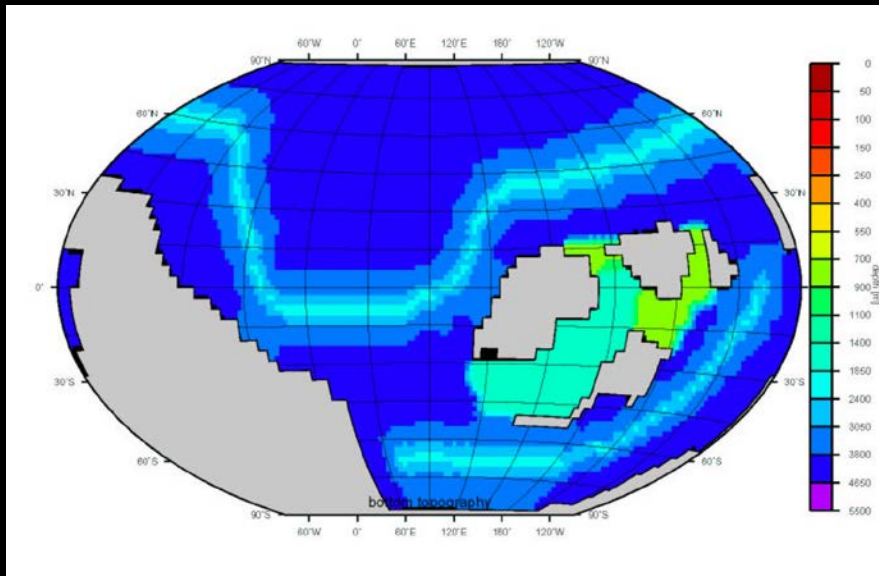
Bathymetry



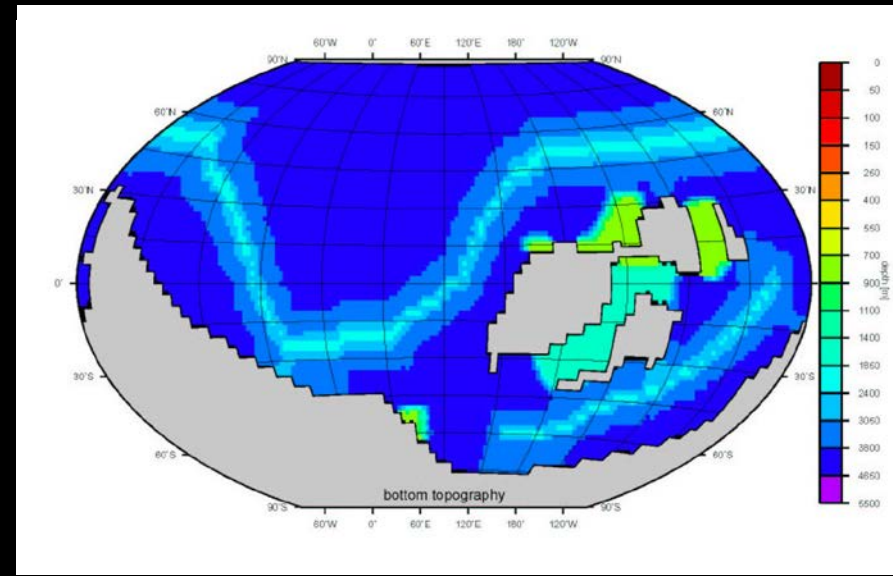
Caradocian High Sea Level



Ashgillian High Sea Level

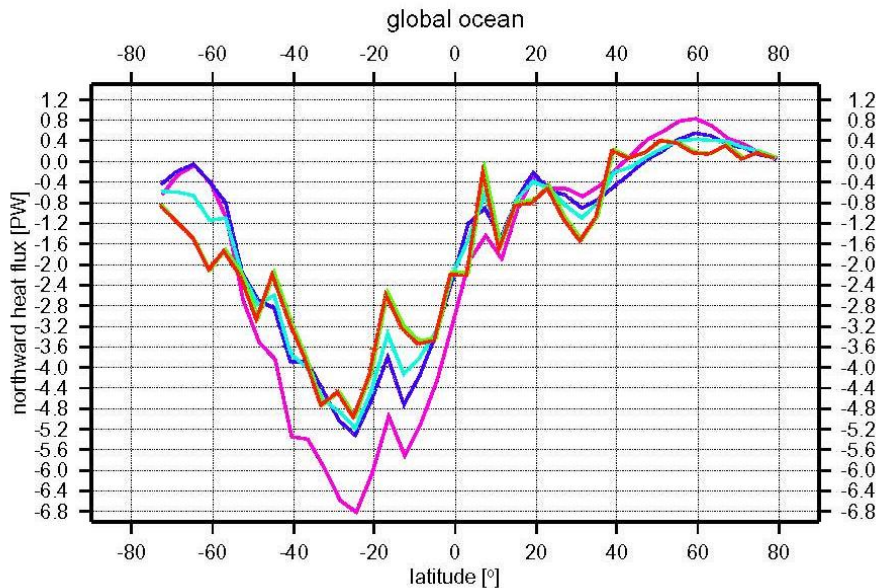


Caradocian Low Sea Level

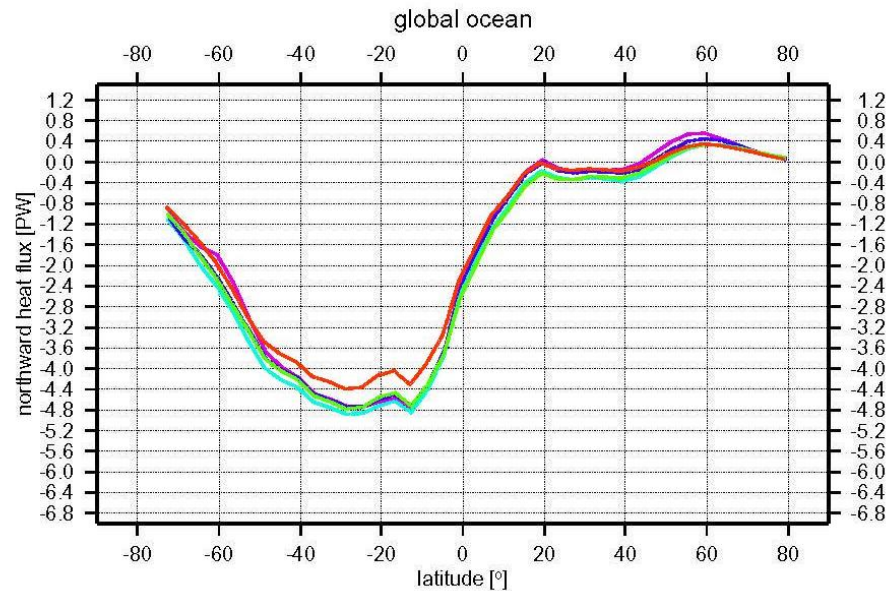


Ashgillian Low Sea Level

Ocean Heat Transport and atmospheric pCO₂



Caradocian

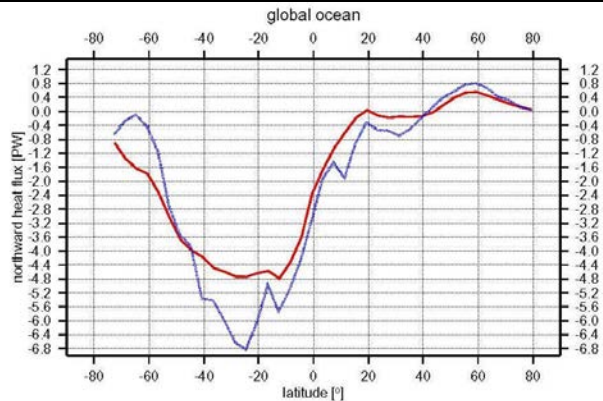


Ashgillian

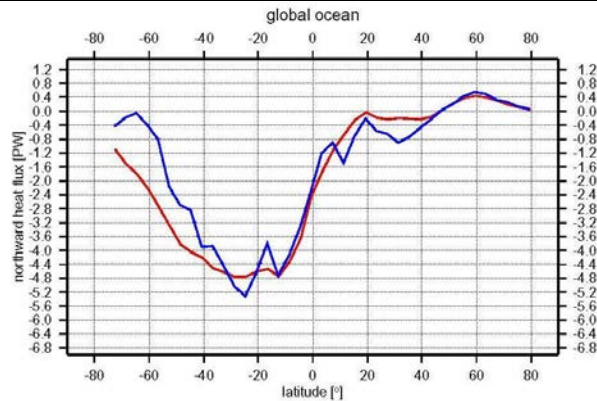
8x PAL 10x PAL 12x PAL 15x PAL 18x PAL

Ocean Heat Transport and Paleogeography

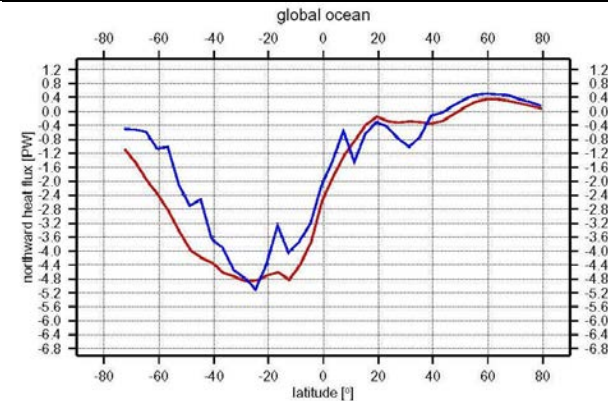
8x PAL



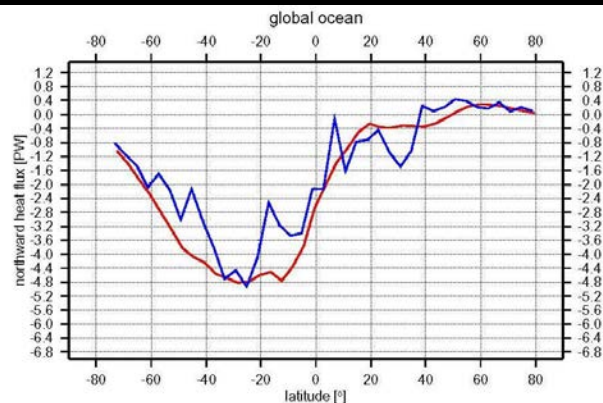
10x PAL



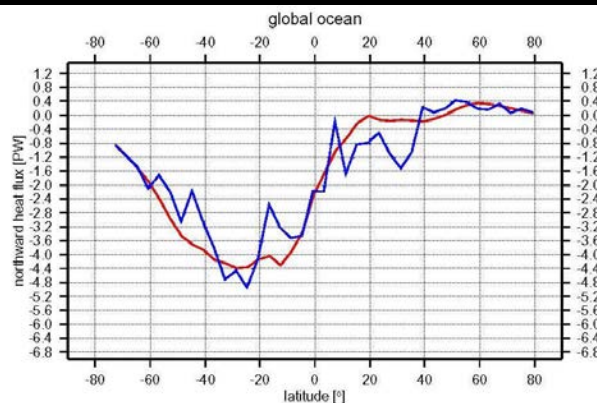
12x PAL



15x PAL



18x PAL



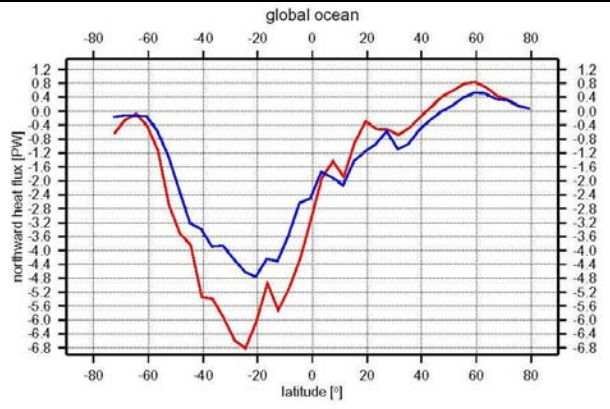
Caradocian



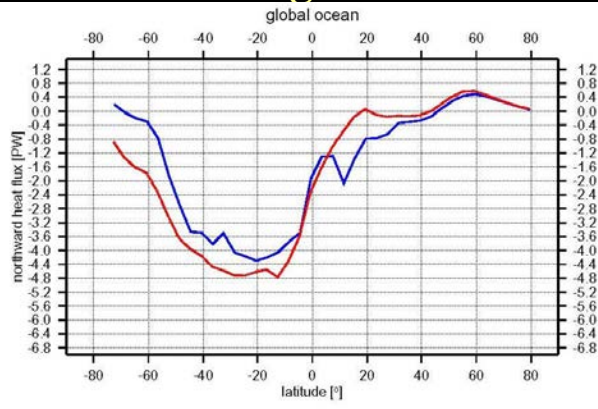
Ashgillian

Ocean Heat Transport and Sea Level

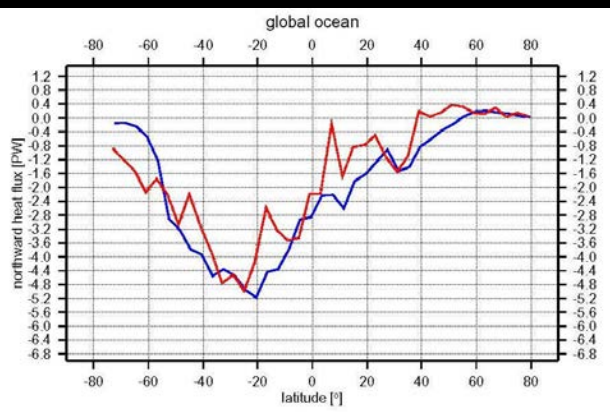
8x Caradocian



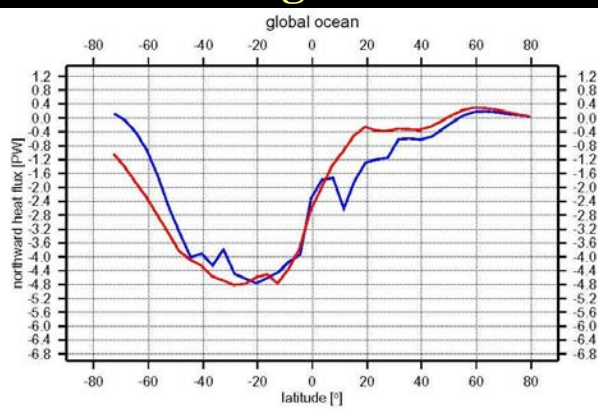
8x Ashgillian



15x Caradocian



15x Ashgillian



 Low Sea Level

 High Sea Level

Main Outcomes - Ocean (1)

Our primary results indicate that:

- in the Caradocian, a decrease in atmospheric $p\text{CO}_2$ leads to an increase in poleward ocean heat transport;
 - in the Ashgillian, there is no change in poleward ocean heat transport between 15x and 8x PAL
- therefore, a decrease in atmospheric $p\text{CO}_2$ levels would not lead to a positive feedback favoring glaciation

Main Outcomes - Ocean (2)

- With atmospheric $p\text{CO}_2$ values above 8x PAL, the Ashgillian paleogeography leads to a higher ocean heat transport in the Southern Hemisphere.
- Only in the simulation with 8x PAL does the paleogeographic change lead to a lower heat transport in lower Southern latitudes during the Ashgillian

Main Outcomes - Ocean (3)

- at 8x and 10x PAL, low sea level simulations have a lower southward ocean heat transport → a drop in sea level in the Ashgillian at low atmospheric pCO₂ levels could have lead to a positive feedback favoring glaciation by reducing the southward ocean heat transport in addition to the ice-albedo feedback of the exposed shelves

Acknowledgements

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